

# Investigating the climate change response of the NH wintertime storm tracks

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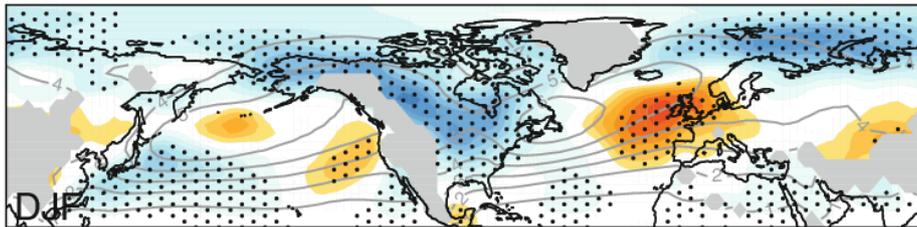
National Centre for Atmospheric Science

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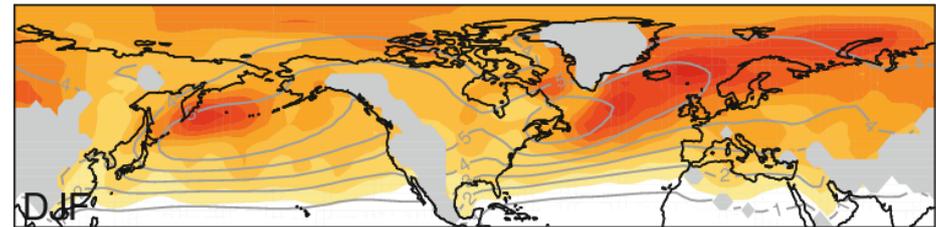
Thanks to Ben Harvey, Tim Woollings, Giuseppe Zappa, Kevin Hodges

1. What governs the mean response?
2. Why is the inter-model spread so large?

DJF CMIP5 Storm Track response

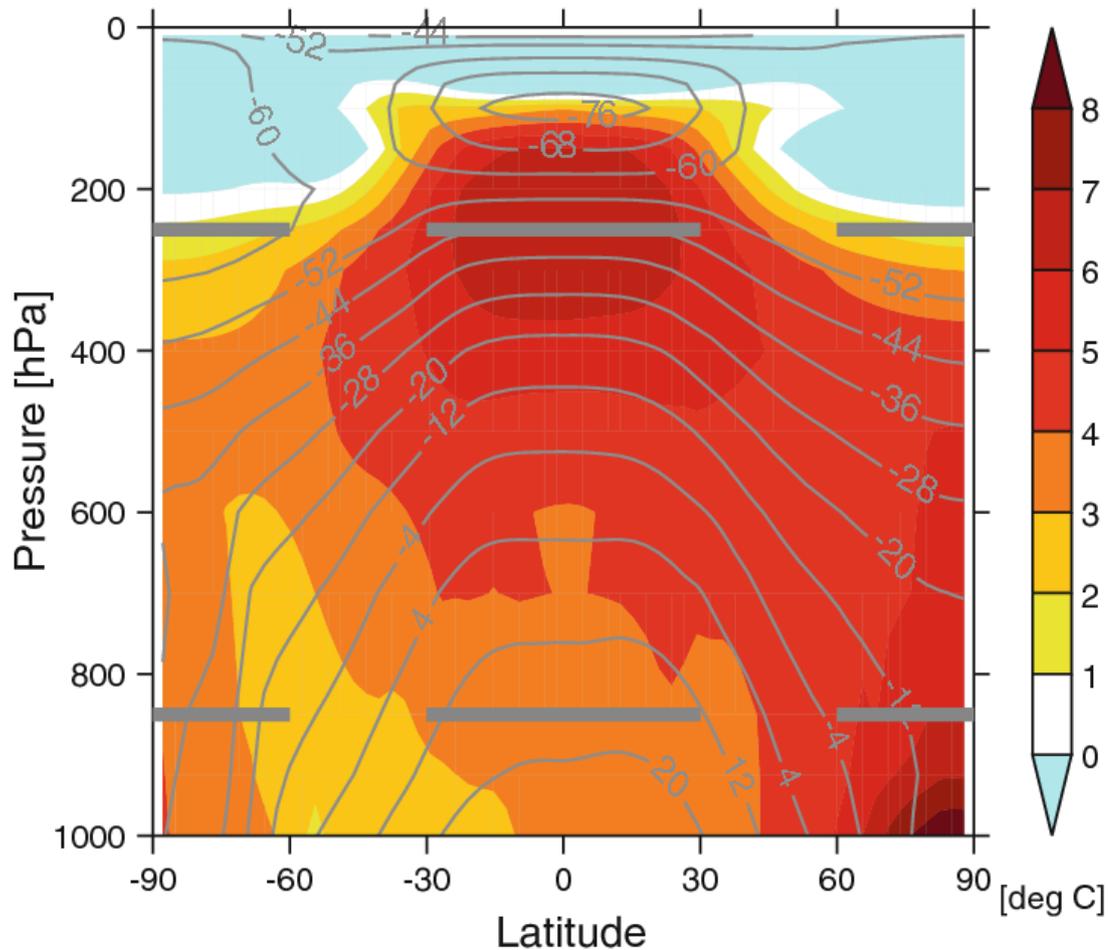


DJF CMIP5 inter-model spread



Left) DJF CMIP5 multi-model mean storm track response (RCP4.5 2080-2100 minus HIST 1985-2005) measured by standard deviation of the 2-6 bandpass filtered mean sea level pressure. Right) DJF CMIP5 inter-model standard deviation of storm track responses





Zonal mean climate change temperature responses from the CMIP5 climate models

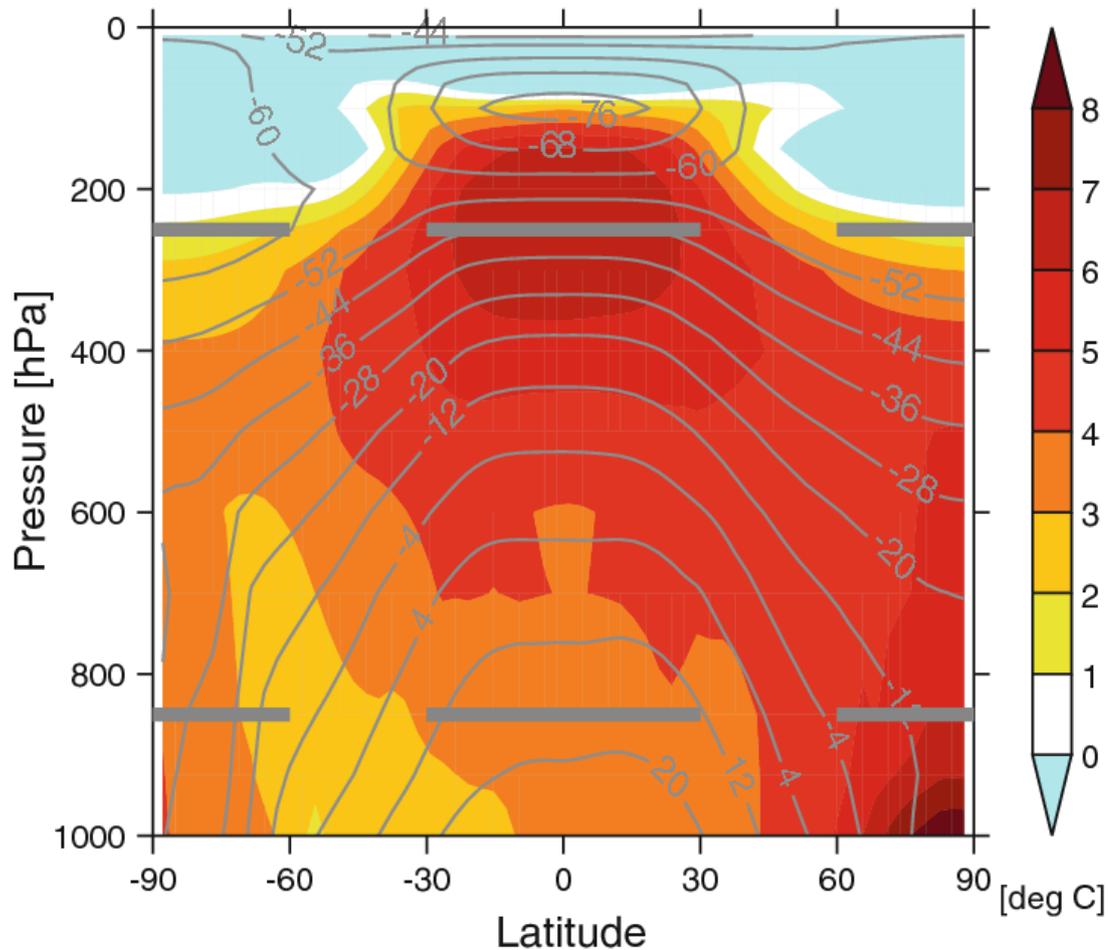
Warming in the:

1. Upper Tropics
2. Lower Polar troposphere

Using an idealised dry GCM, Butler et al. (2010) found that:

Upper Tropics heating = Intensified (and poleward shifted) storm track

Lower Polar heating = Weakened (and equatorward shifted) storm track



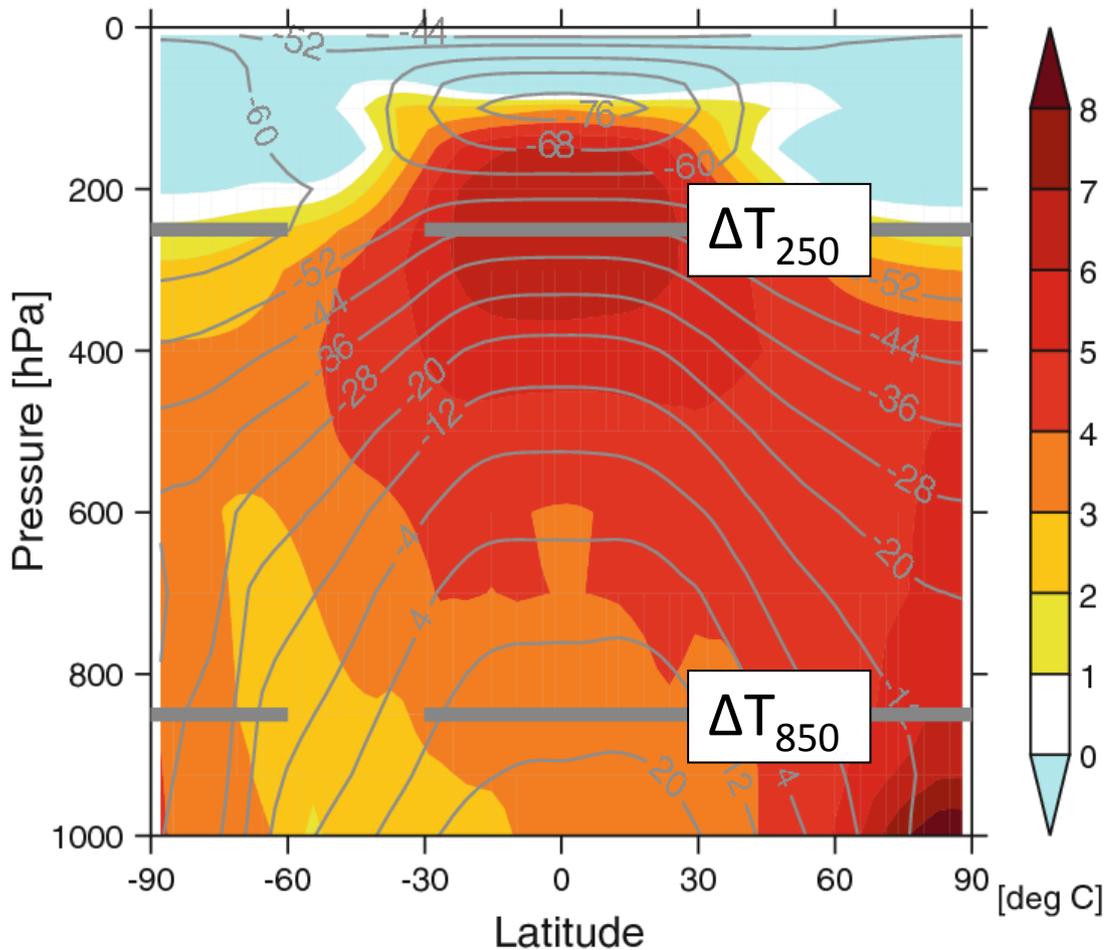
Zonal mean climate change temperature responses from the CMIP5 climate models

Warming in the:

1. Upper Tropics
2. Lower Polar troposphere

Investigate this:

1. In the CMIP5 climate model ensemble (Harvey et al. 2013)
2. By performing climate model sensitivity experiments (Harvey et al. 2015)



Zonal mean climate change temperature responses from the CMIP5 climate models

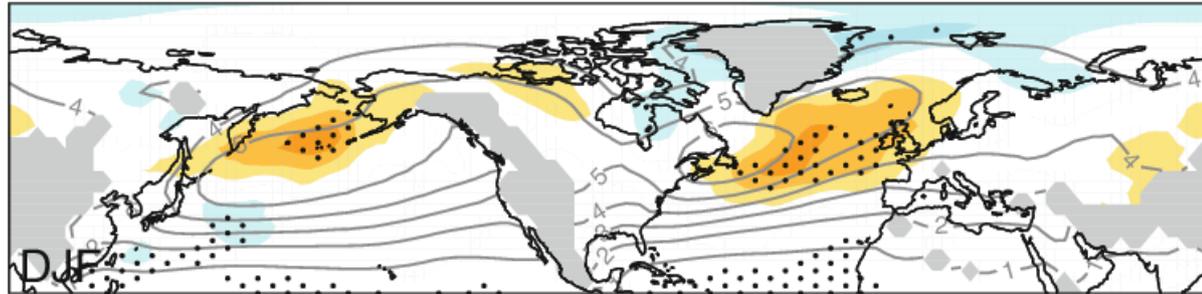
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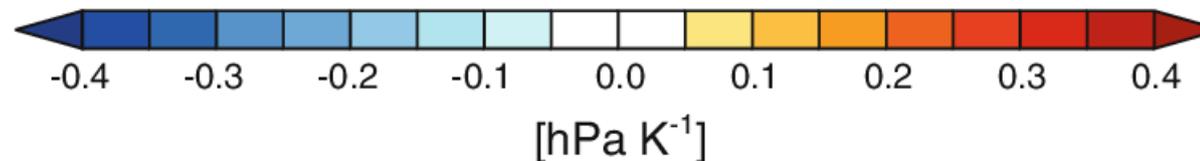
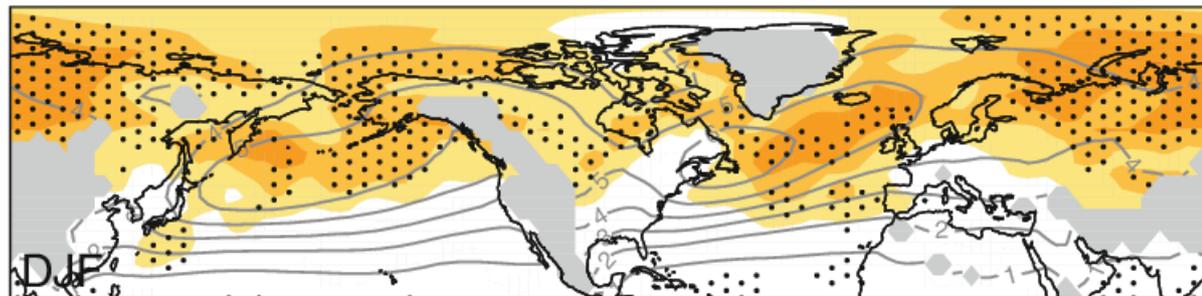
Investigate this:

1. In the CMIP5 climate model ensemble (Harvey et al. 2013)
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(b) winter  $\Delta T_{250}$  regression slope



(a) winter  $\Delta T_{850}$  regression slope

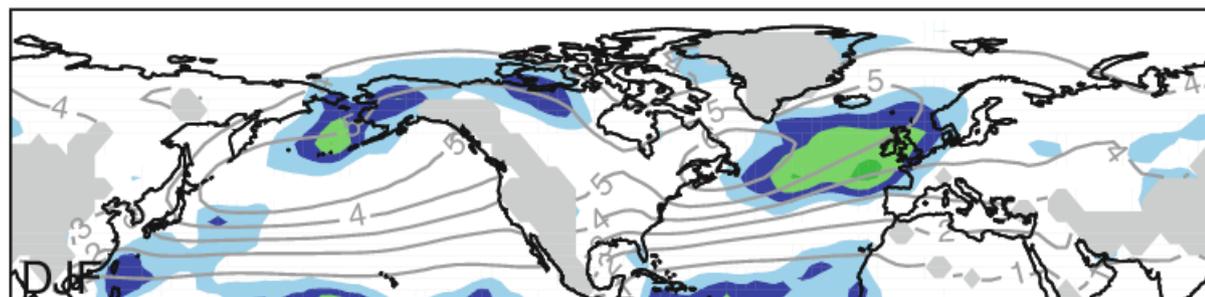


$$\Delta T_{850} = T_{\text{TROPICS}}_{850} \text{ minus } T_{\text{POLAR}}_{850}$$

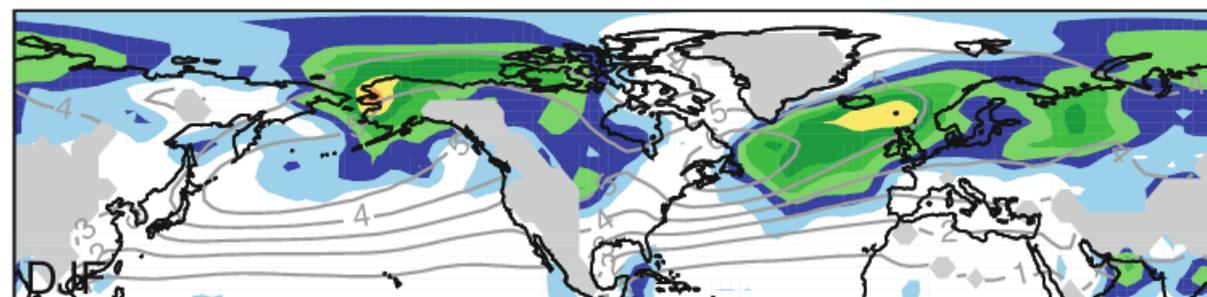
Regression coefficients for DJF CMIP5 storm track responses against lower) the 850hPa temperature differences and upper) the 250hPa temperature differences.

Both  $\Delta T_{250}$  and  $\Delta T_{850}$  appear to be important for the mean response of the NH DJF storm tracks

$\Delta T$  Atlantic 250hPA



$\Delta T$  Atlantic 850hPA

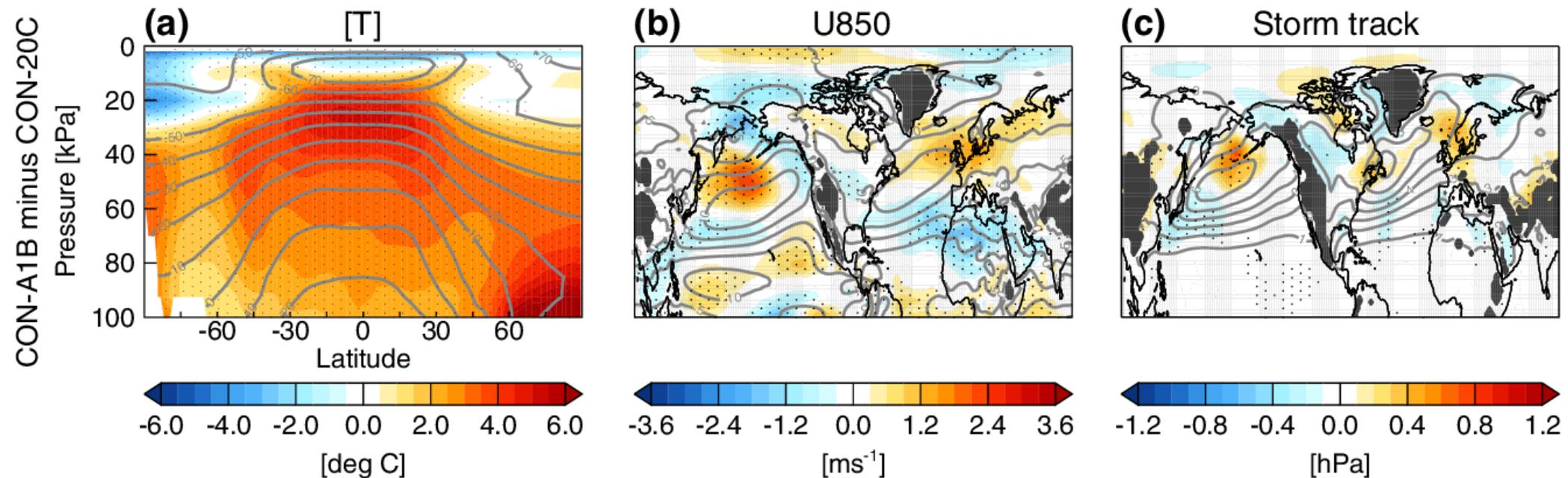


Fraction of the inter-model spread in the DJF CMIP5 storm track responses explained by (lower) linear regression with equator-to-pole 850hPa Atlantic temperature differences and (upper) with the 250hPa Atlantic temperature differences.

The inter-model spread in CMIP5 North Atlantic storm track responses is more associated with Atlantic  $\Delta T_{850}$

Perform sensitivity experiments with the N96 L38 HadGAM1 atmosphere model with forcings derived from the CMIP3 ensemble:

1. **Present day experiment** with multi-model mean CMIP3 SST and constructed sea ice
2. **Climate change experiment** (A1B scenario) with multi-model mean CMIP3 SST and constructed sea ice





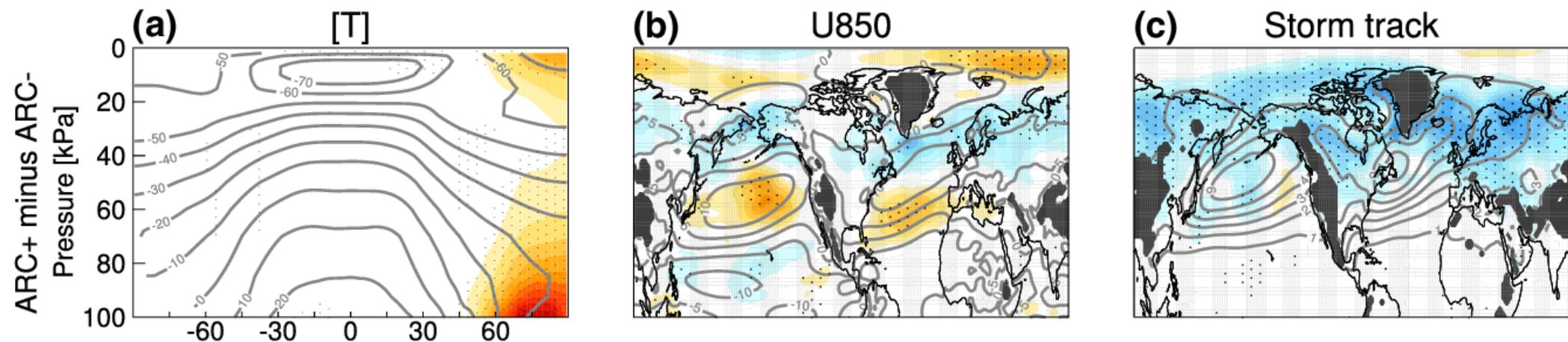
Perform sensitivity experiments with the N96 L38 HadGAM1 atmosphere model with forcing from CMIP3 ensemble:

1. **Present day experiment** with multi-model mean CMIP3 SST and constructed sea ice
2. **Climate change experiment** (A1B scenario) with multi-model mean CMIP3 SST and constructed sea ice

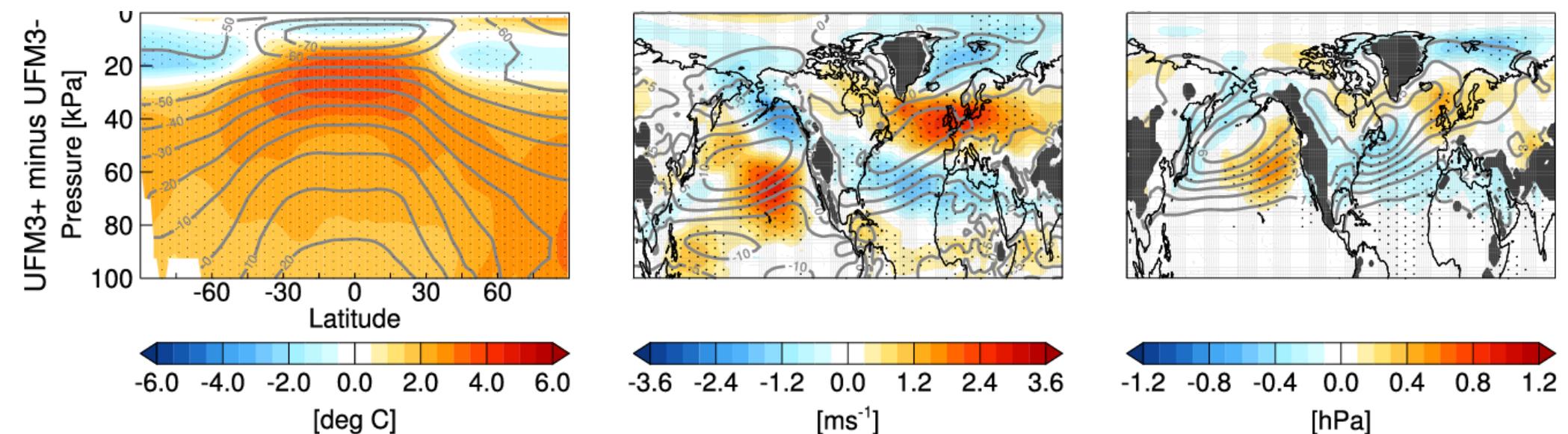
Two sets of sensitivity experiments intended to simulate the CMIP3 inter-model spread (plus and minus one standard deviation) :

3. **Large Arctic Warming** and **Small Arctic Warming** by applying Arctic sea ice perturbations and high-latitude SST perturbations
4. **Large Tropical Warming** and **Small Tropical Warming** by applying globally uniform SST perturbations

## Warm Arctic minus Cold Arctic



## Warm Tropics minus Cold Tropics



1. Analysis of the CMIP5 ensemble and sensitivity experiments with HadGEM1 suggest that
  - Upper Tropical Warming intensifies the NH DJF storms tracks
  - Lower Polar Warming weakens the NH DJF storm tracks
2. Both Upper Tropical and Lower Polar Warming are important for the mean response, but the Lower Polar Warming is more important for the large spread of projections in the North Atlantic (see also Son et al. 2015).

## Caveats:

1. Have we chosen the correct processes to look at – e.g. role of SST changes in the North Atlantic (Woollings et al. 2012) or stratospheric responses (Manzini et al. 2014)?
2. The sensitivity experiments are with only one model – is robustness an issue? Are there better ways of performing the experiments?